

Modeling and Simulation in Rotationally Constrained (Convective) Flows

Keith Julien¹

Ian Grooms^{4,1}, Antonio Rubio¹, Geoff Vasil³,
Baylor Fox-Kemper⁶, Edgar Knobloch², Jeff Weiss¹,
Michael Calkins¹, Philippe Marti¹, Jon Aurnou⁵

¹ Department of Applied Mathematics, University of Colorado Boulder

² Department of Physics, University of California Berkeley

³ Department of Mathematics, Sydney

⁴ Courant Institute of Mathematical Sciences, New York University

⁵ Earth Sciences, UCLA

⁶ Geological Sciences, Brown University

Support: NSF FRG DMS
NSF EAR CSEDI

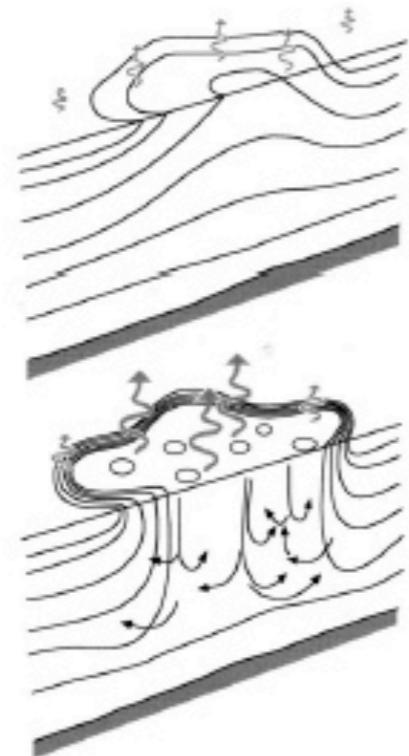
Outline

- Motivation
- Slow Manifold Equations: Nonhydrostatic QG limit
- Application - Simulations

Rotationally Constrained Convective Flows in GAFD $Ro \ll 1$

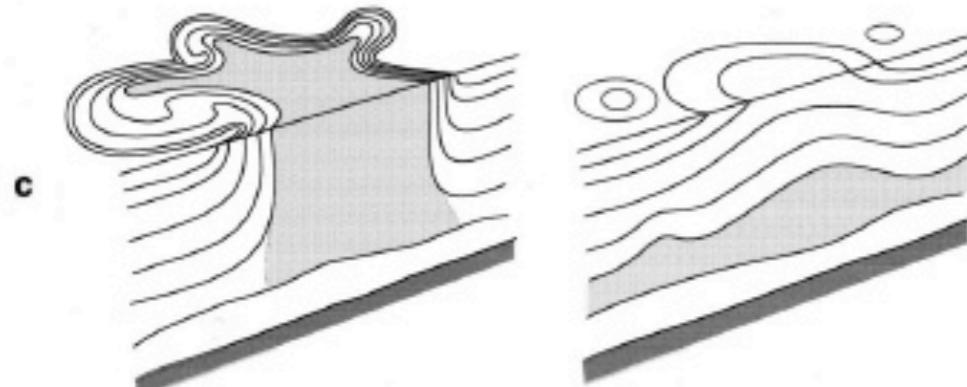
Marshall and Schott: OPEN-OCEAN CONVECTION • 5

Preconditioning
~100 Km patch



$$\begin{aligned} Ro &\sim 0.1 - 0.4 \\ U &\sim 0.05 \text{ m/s} \\ \Omega &\sim 7 \times 10^{-5} \text{ rad/s} \\ L &\sim 2 \text{ km} \end{aligned}$$

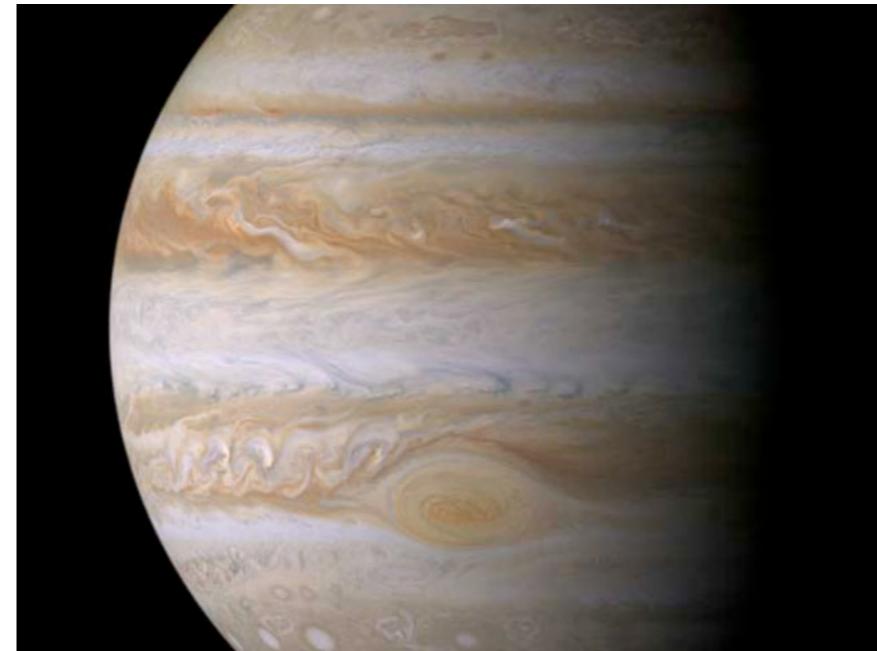
Deep b
Convection



Lateral exchange & spreading

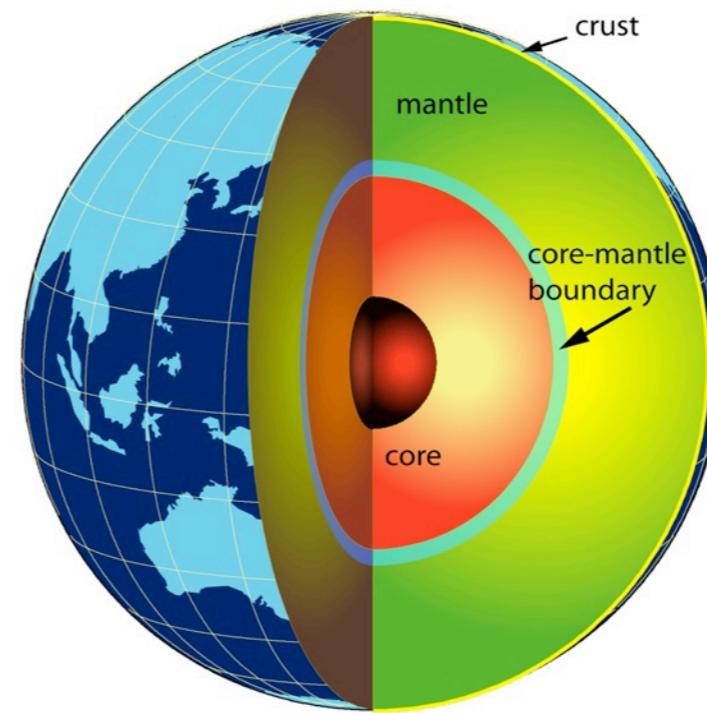
$$Ro = \frac{t_\Omega}{t_{adv}} = \frac{U}{2\Omega L}$$

large-scale flow generation on Giant Planets



$$\begin{aligned} Ro &\sim 10^{-2} \\ U &\sim 100 \text{ m/s} \\ \Omega &\sim 2 \times 10^{-4} \text{ rad/s} \\ L &\sim 15 \text{ Mm} \end{aligned}$$

turbulence primary driver for geomagnetic field

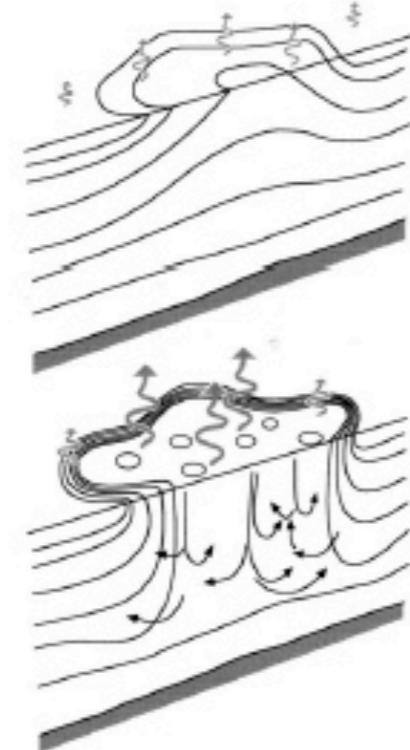


$$\begin{aligned} Ro &\sim 10^{-7} \\ U &\sim 3 \times 10^{-4} \text{ m/s} \\ \Omega &\sim 7 \times 10^{-5} \text{ rad/s} \\ L &\sim 2260 \text{ km} \end{aligned}$$

Rotationally Constrained Convective Flows in GAFD $Ro \ll 1$

Marshall and Schott: OPEN-OCEAN CONVECTION • 5

Preconditioning

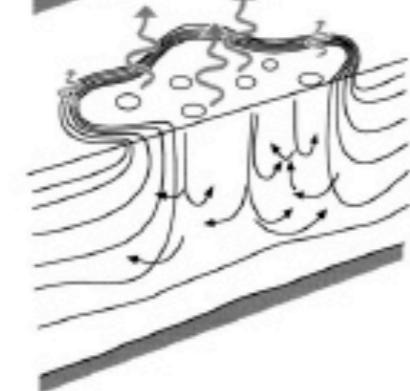


$$Ro \sim 0.1 - 0.4$$

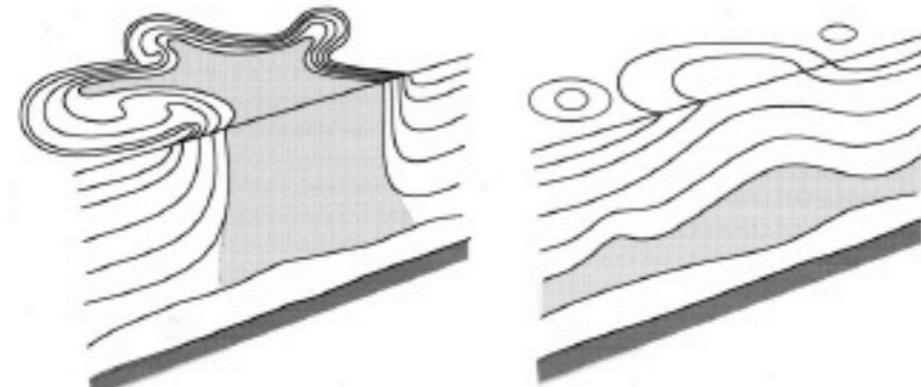
$$Re \sim 10^8$$

$$Ek \sim 10^{-9}$$

Deep Convection



c



Lateral exchange & spreading

$$Ro = \frac{t_\Omega}{t_{adv}} = \frac{U}{2\Omega L}$$

large-scale flow generation on Giant Planets

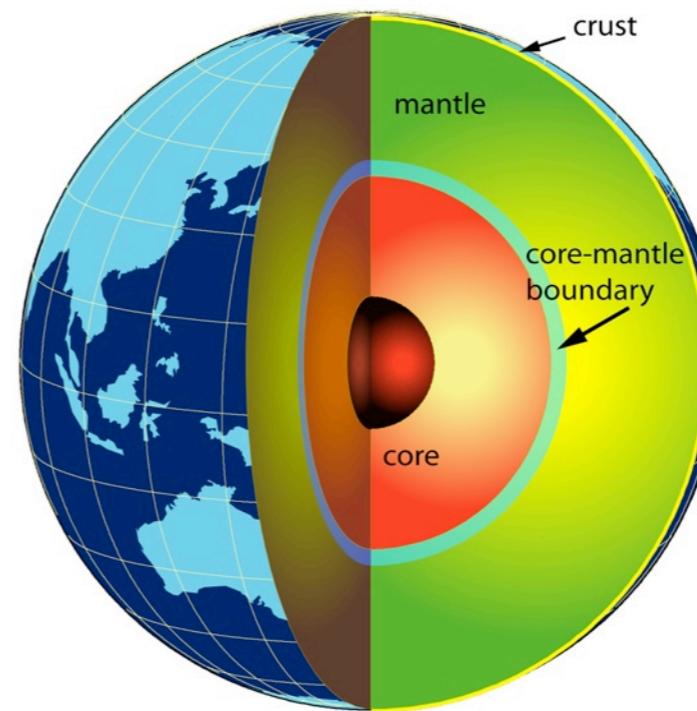


$$Ro \sim 10^{-2}$$

$$Re \sim 10^{16}$$

$$Ek \sim 10^{-18}$$

turbulence primary driver for geomagnetic field



$$Ro \sim 10^{-7}$$

$$Re \sim 10^8$$

$$Ek \sim 10^{-15}$$

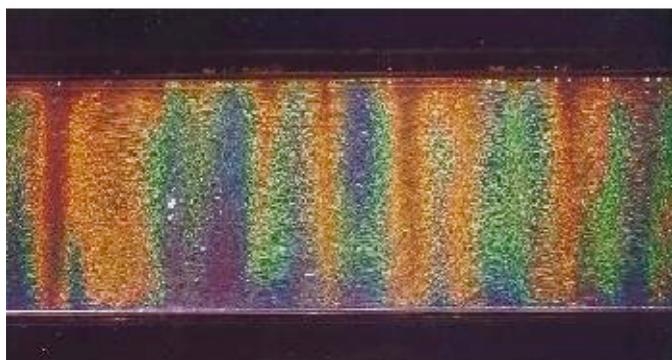
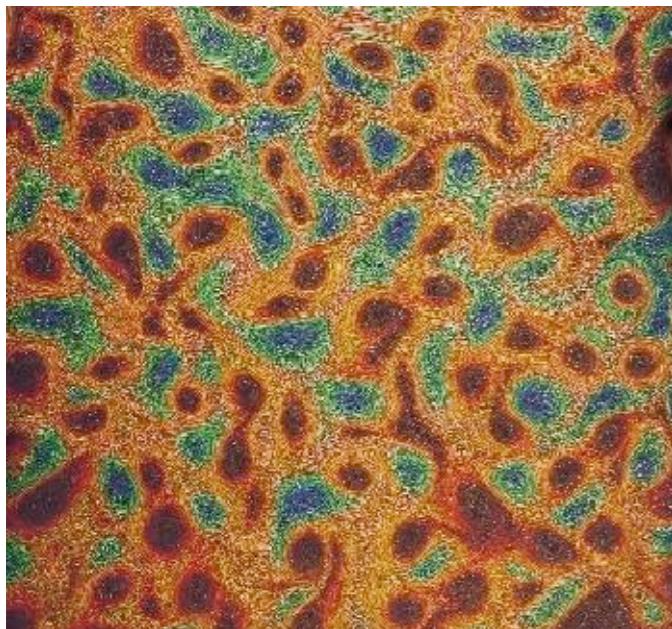
Nonhydrostatic Investigations: Canonical Configurations

Maxworthy & Narimosa JPO 1994

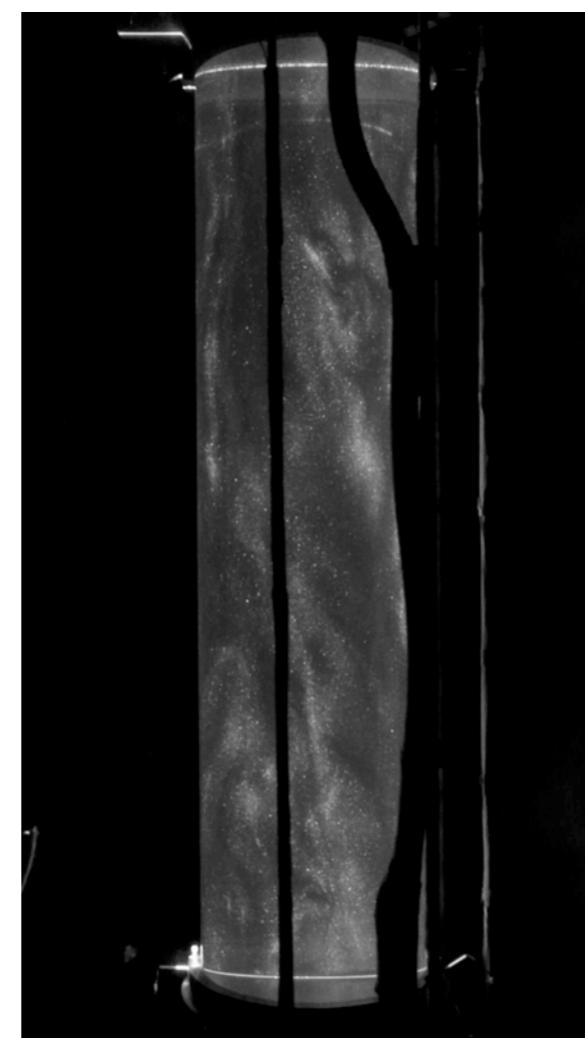
- Non-hydrostatic, rotationally constrained flows characterized by columnar structures
- probing low Ro, high Re challenging
 - experimentation: restricted by mechanical and fluid properties
 - DNS: restricted by spatiotemporal resolution constraints.

Liu & Ecke PRE '09; King et al Nature '09

Sakai, JFM 1997: $\text{RaE}^{4/3} = 36$, $\text{Ro} \approx 0.1$, $\sigma = 7$

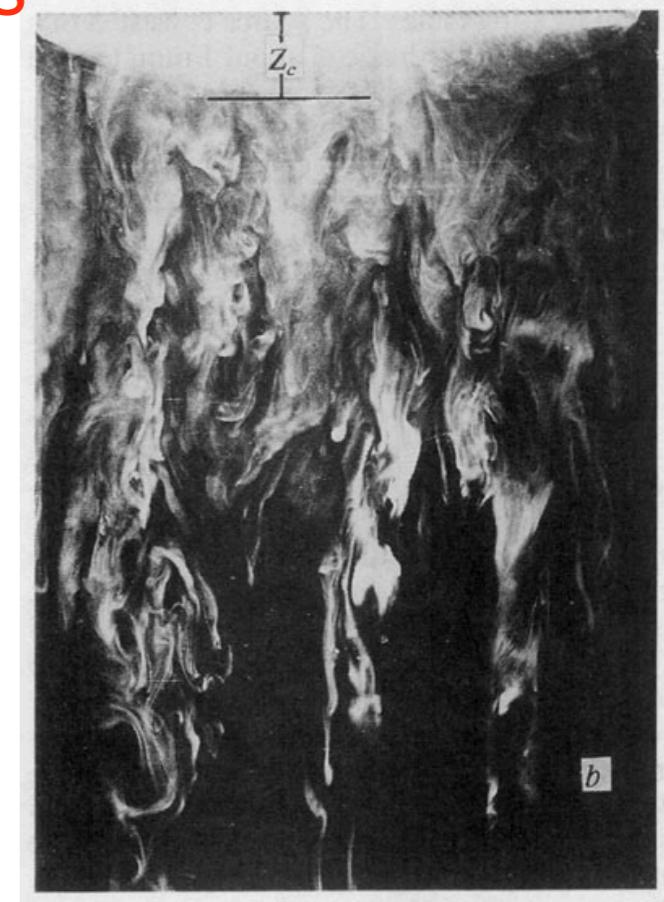


Taylor columns

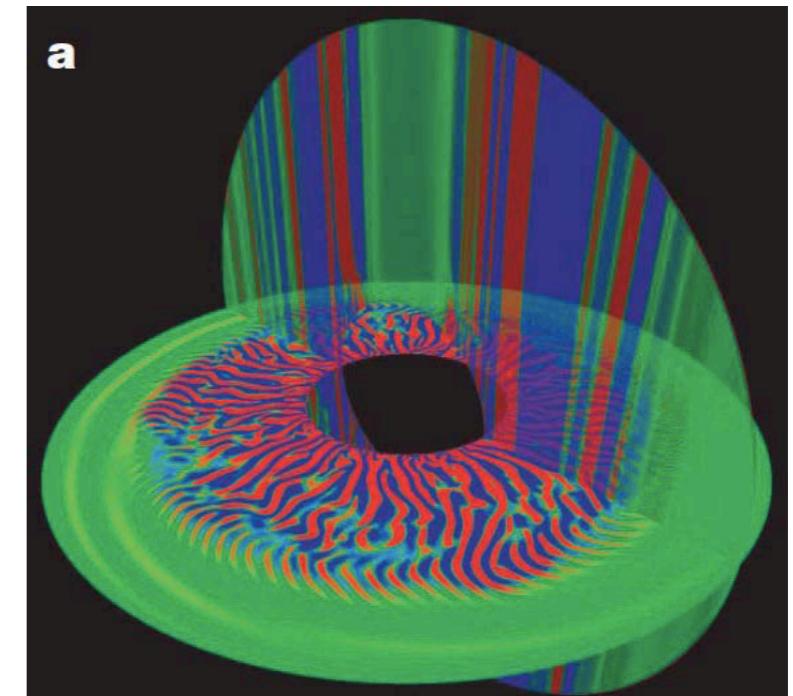


Plumes

Aurnou, $\text{RaE}^{4/3} = 755$, $\text{Ro} \approx 0.13$, $\sigma = 7$



nonhomog. heat source



Axial vorticity, $E \sim 2e-7$ (Kageyama et al Nature 2008)
 $\text{Ra} = 1.5e10$

Navier-Stokes Equations (Non-dimensional Characterization)

- Generic non-dimensionalization: $L, U, \Delta T, P$

$$D_t \mathbf{u} + Ro^{-1} \hat{\mathbf{z}} \times \mathbf{u} = -Eu \nabla p + \Gamma b \hat{\mathbf{z}} + Re^{-1} \nabla^2 \mathbf{u} + \mathbf{S}$$

$$D_t b - \Gamma^{-1} Fr^{-2} w \partial_z \rho(z) = Pe^{-1} \nabla^2 b$$

$$\nabla \cdot \mathbf{u} = 0$$

where $D_t := \partial_t + \mathbf{u} \cdot \nabla$ with (\mathbf{u}, p, b) for velocity, pressure & buoyancy fields.

- Non-dimensional Parameters:

Rossby Number

$$Ro = \frac{U}{2\Omega L}$$

Euler Number

$$Eu = \frac{P}{\rho_0 U^2}$$

Buoyancy Number

$$\Gamma = \frac{g\alpha\Delta T L}{U^2}$$

Froude Number

$$Fr = \frac{U}{NL}$$

Ekman Number

$$Ek = \frac{Ro}{Re} = \frac{\nu}{2\Omega L^2}$$

Reynolds Number

$$Re = \frac{UL}{\nu}$$

Péclet Number

$$Pe = \frac{UL}{\kappa}$$

Navier-Stokes Equations (Incompressible Fluid)

- Generic non-dimensionalization: $L, U, \Delta T, P$

$$D_t \mathbf{u} + Ro^{-1} \hat{\mathbf{z}} \times \mathbf{u} = -Eu \nabla p + \Gamma b \hat{\mathbf{z}} + Re^{-1} \nabla^2 \mathbf{u} + \mathbf{S}$$

$$D_t b - \Gamma^{-1} Fr^{-2} w \partial_z \rho(z) = Pe^{-1} \nabla^2 b$$

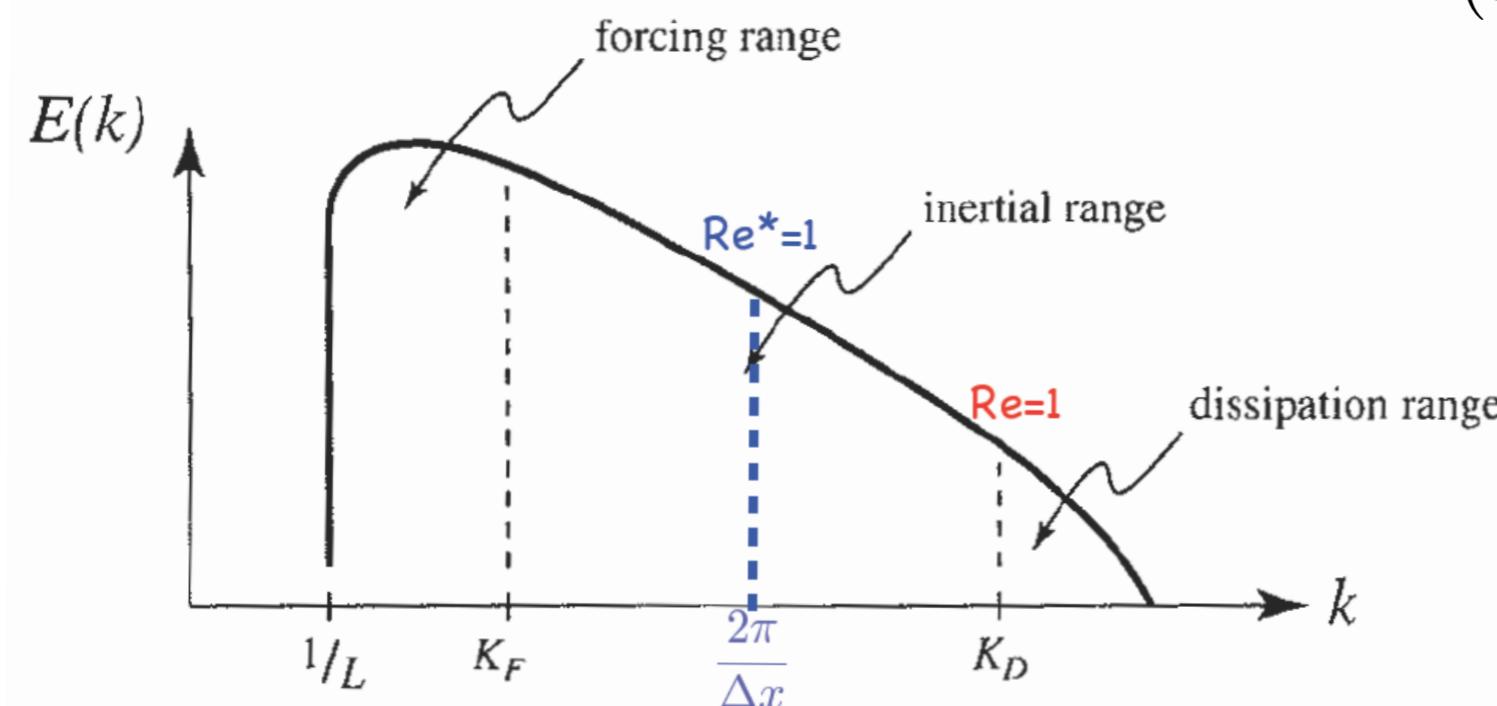
$$\nabla \cdot \mathbf{u} = 0$$

where $D_t := \partial_t + \mathbf{u} \cdot \nabla$ with (\mathbf{u}, p, b) for velocity, pressure & buoyancy fields.

- Turbulence Challenge: d.o.f. (grid-pts/modes) $\Rightarrow N^3 \sim Re^{\frac{9}{4}}$ (Pope, 2000)

$$(10^{6+})^3 \sim (10^{8+})^{\frac{9}{4}} \Rightarrow \text{GAFD}$$

$$(10^3)^3 \sim (10^4)^{\frac{9}{4}} \Rightarrow \text{DNS}$$



$$\mathcal{T} \sim 2Re^3 / P_{flop \ rate}$$

$$30d \sim 2(10^8)^3 / 10^{23} \Rightarrow \text{GAFD}$$

Moore's Law $\Rightarrow 70$ yrs away

Navier-Stokes Equations (Incompressible Fluid)

- Generic non-dimensionalization: $L, U, \Delta T, P$

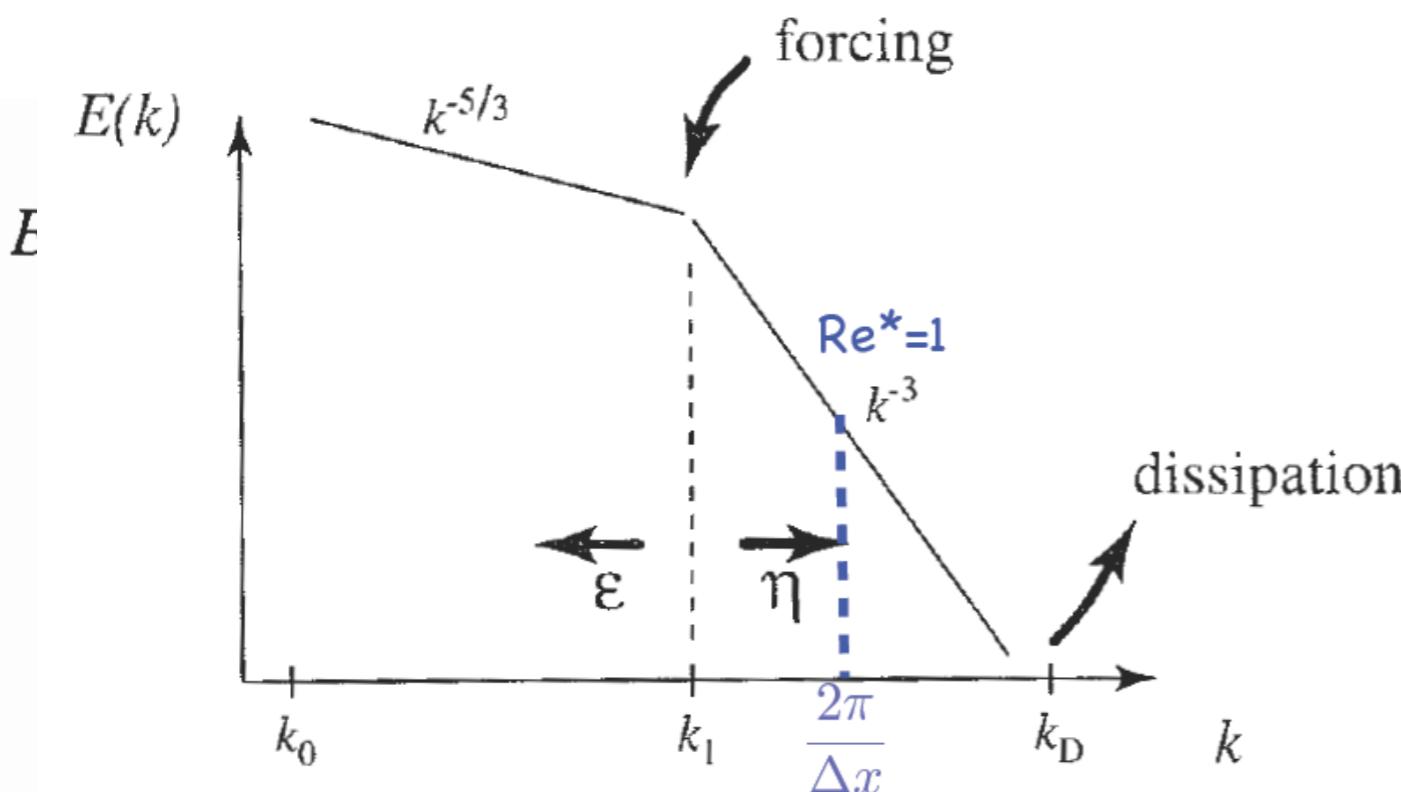
$$D_t \mathbf{u} + Ro^{-1} \hat{\mathbf{z}} \times \mathbf{u} = -Eu \nabla p + \Gamma b \hat{\mathbf{z}} + Re^{-1} \nabla^2 \mathbf{u} + \mathbf{S}$$

$$D_t b - \Gamma^{-1} Fr^{-2} w \partial_z \rho(z) = Pe^{-1} \nabla^2 b$$

$$\nabla \cdot \mathbf{u} = 0$$

where $D_t := \partial_t + \mathbf{u} \cdot \nabla$ with (\mathbf{u}, p, b) for velocity, pressure & buoyancy fields.

- Turbulence Challenge: d.o.f. (grid-pts/modes) $\Rightarrow N^3 \sim Re^{\frac{9}{4}}$ (Pope, 2000)



$$(10^{6+})^3 \sim (10^{8+})^{\frac{9}{4}} \Rightarrow \text{GAFD}$$

$$(10^3)^3 \sim (10^4)^{\frac{9}{4}} \Rightarrow \text{DNS}$$

$$\mathcal{T} \sim 2Re^3 / P_{flop \ rate}$$

$$30d \sim 2(10^8)^3 / 10^{23} \Rightarrow \text{GAFD}$$

Moore's Law \Rightarrow 70 yrs away

Navier-Stokes Equations: Rotationally Constrained Flows, $Ro \ll 1$

- For $Ro \ll 1$ turbulence challenge compounded

$$\boxed{D_t \mathbf{u} + Ro^{-1} \hat{\mathbf{z}} \times \mathbf{u} = -Eu \nabla p + \Gamma b \hat{\mathbf{z}} + Re^{-1} \nabla^2 \mathbf{u}}$$

$$D_t b - \Gamma^{-1} Fr^{-2} w \partial_z \rho(z) = Pe^{-1} \nabla^2 b$$

$$\boxed{\nabla \cdot \mathbf{u} = 0}$$

- NSE a stiff PDE, \exists fast inertial waves & slow geostrophically balanced eddies

Fast Inertial Waves

$$\omega_{fast} \sim Ro^{-1} \frac{k_z}{\sqrt{k_\perp^2 + k_z^2}}$$

of secondary importance

Geostrophic Eddies/Slow Waves

$$\omega_{slow} \sim \mathcal{O}(1)$$

$$Ro^{-1} \hat{\mathbf{z}} \times \mathbf{u} \approx -Eu \nabla p, \quad \nabla \cdot \mathbf{u} = 0 \Rightarrow$$

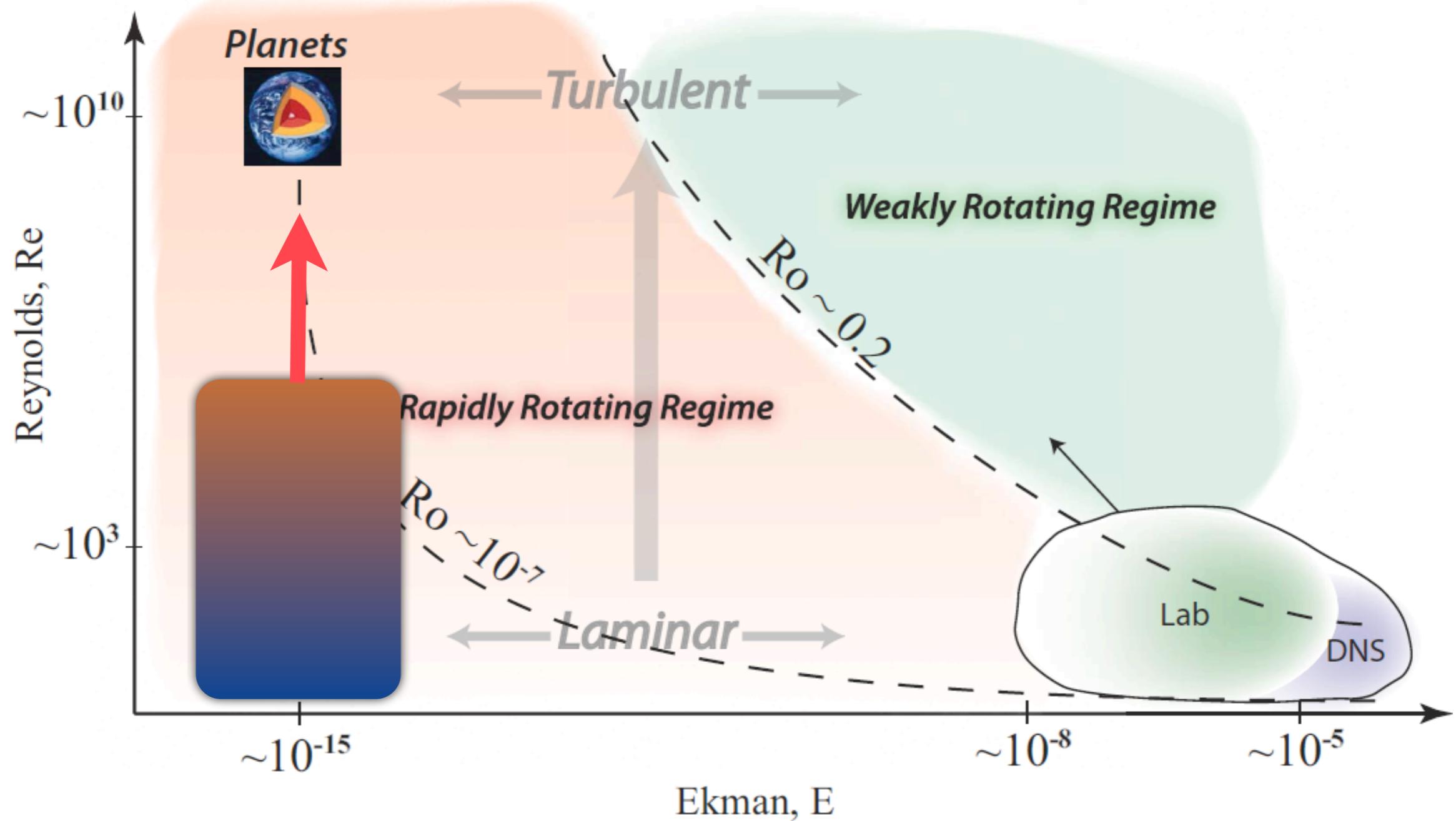
$$\hat{\mathbf{z}} \cdot \nabla(\mathbf{u}, p) \approx 0$$

Proudman-Taylor Thm (1916,1923)

Low Rossby Number Challenge

- Fast waves + geostrophically balanced eddies limit DNS/Lab investigations

Resolution: Quasi-Geostrophic Theory. Restrict dynamics to geostrophic manifold and identify Reduced (Nonhydrostatic) PDE's!



Reduced QG Equations: Perturbation Theory

- Select aspect ratio of interest, set distinguished limits
- Perform asymptotic expansion in Rossby number, $\text{Ro} \ll 1$

$$u = u_0 + \text{Ro} u_1 + \text{Ro}^2 u_2 + \dots$$

$$v = v_0 + \text{Ro} v_1 + \text{Ro}^2 v_2 + \dots$$

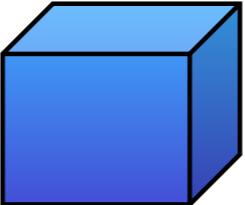
⋮
⋮

- Projection to slow manifold J. et al JFM 06, J. & Knobloch JMP 07, Calkins et al JFM 13
- Solve sequence of LPDE's with secularity conditions

Rotationally constrained flows and aspect ratio

$Ro \ll 1$

$A = H/L$

QG	Intermediate	Convection
 $H/L \ll 1$ Charney (1948)	 $H/L = O(1)$ Embid & Majda (1998)	 $H/L \gg 1$ J. et al TCFD '98; JFM 06

- **Unified QG approach:**

pointwise geostrophy: $Ro^{-1} \hat{z} \times \mathbf{u}_\perp = -E u \nabla_\perp p$

inc. vortex stretching:

$$\frac{U^*}{L} \cdot \frac{U^*}{L} \sim \frac{2\Omega W^*}{H}$$

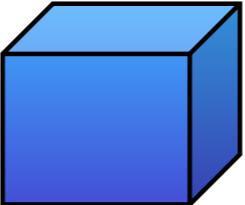
vert. velocity scaling:

$$w_0 = \mathcal{O}(ARo)$$

Rotationally constrained flows and aspect ratio

$Ro \ll 1$

$A = H/L$

QG	Intermediate	Convection
 $H/L \ll 1$ Charney (1948)	 $H/L = O(1)$ Embid & Majda (1998)	 $H/L \gg 1$ J. et al TCFD '98; JFM 06

- **Unified QG approach:** $\partial_z \rightarrow A^{-1} \partial_z, \quad w = ARoW$

$$\hat{\mathbf{z}} \times \mathbf{u}_\perp = -\nabla_\perp p, \quad p = \psi$$

$$D_t^\perp \nabla_\perp^2 \psi - \partial_z W = Re^{-1} \nabla_A^2 \nabla_\perp^2 \psi$$

$$(ARo)^2 D_t^\perp W = (-\partial_z \psi + b) + (ARo)^2 Re^{-1} \nabla_A^2 W$$

$$D_t^\perp b - W \partial_z \rho(z) = Pe^{-1} \nabla_A^2 b$$

- **Unified distinguished limits:** $Eu = Ro^{-1}, \quad \Gamma = (ARo)^{-1}, \quad Fr = ARo$

H-QG

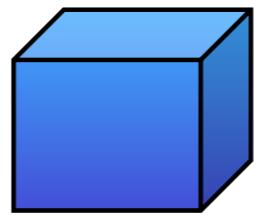
$Ro \ll 1$

$A = H/L$

QG



Intermediate



$H/L \ll 1$

Charney (1948)

$H/L = O(1)$

Embid & Majda (1998)

Conservation: Energy, Enstrophy

PV is the dynamical variable

- **Unified QG approach:** $\partial_z \rightarrow A^{-1} \partial_z, w = ARoW$

$$\hat{\mathbf{z}} \times \mathbf{u}_\perp = -\nabla_\perp p, \quad p = \psi$$

$$D_t^\perp \nabla_\perp^2 \psi - \partial_z W = Re^{-1} \nabla_A^2 \nabla_\perp^2 \psi$$

$$0 = (-\partial_z \psi + b)$$

$$D_t^\perp b - W \partial_z \rho(z) = Pe^{-1} \nabla_A^2 b$$

- **Unified distinguished limits:** $Eu = Ro^{-1}, \Gamma = (ARo)^{-1}, Fr = ARo$

NH-QG

Conservation: Energy

PV not a dynamical variable

$$Ro \ll 1 \quad A = H/L$$

Convection



$$H/L \gg 1$$

J. et al TCFD '98; JFM 06

- **Unified QG approach:** $\partial_z \rightarrow A^{-1} \partial_z, \quad w = ARoW$

$$\hat{\mathbf{z}} \times \mathbf{u}_\perp = -\nabla_\perp p, \quad p = \psi$$

$$D_t^\perp \nabla_\perp^2 \psi - \partial_z W = Re^{-1} \nabla_A^2 \nabla_\perp^2 \psi$$

$$(ARo)^2 D_t^\perp W = (-\partial_z \psi + b) + (ARo)^2 Re^{-1} \nabla_A^2 W$$

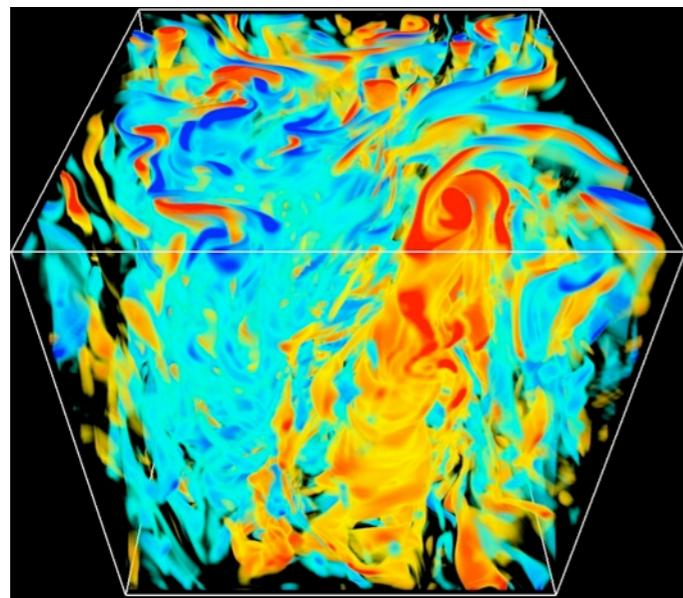
$$D_t^\perp b - W \partial_z \rho(z) = Pe^{-1} \nabla_A^2 b$$

- **Unified distinguished limits:** $Eu = Ro^{-1}, \quad \Gamma = (ARo)^{-1}, \quad Fr = ARo$

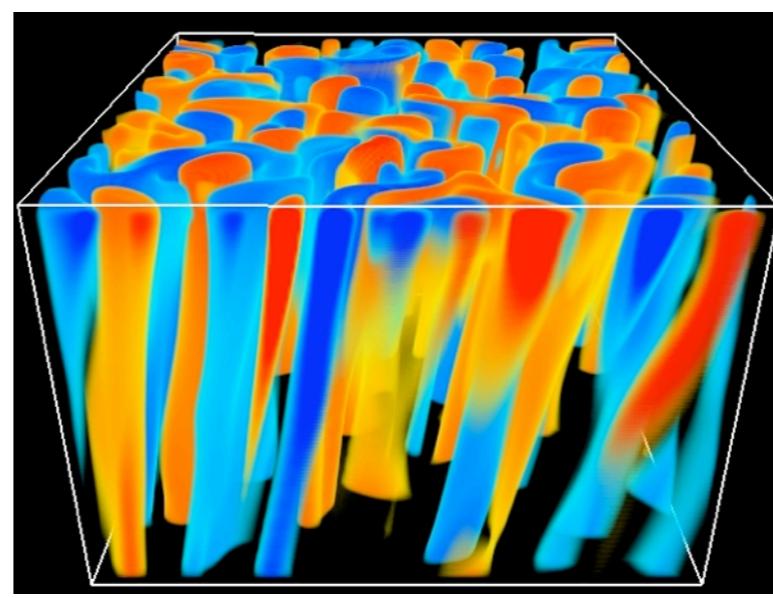
Application to Thermal Convection

Plane layer
convection:

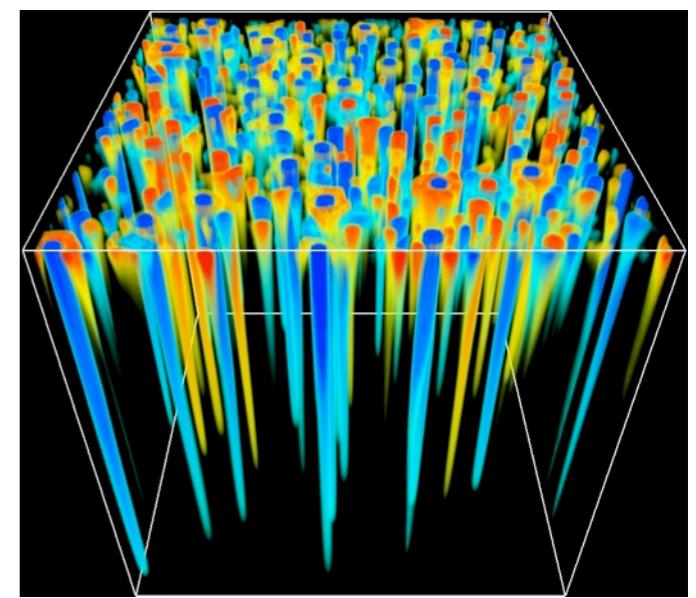
upright turbulent



tilted f-plane, columnar

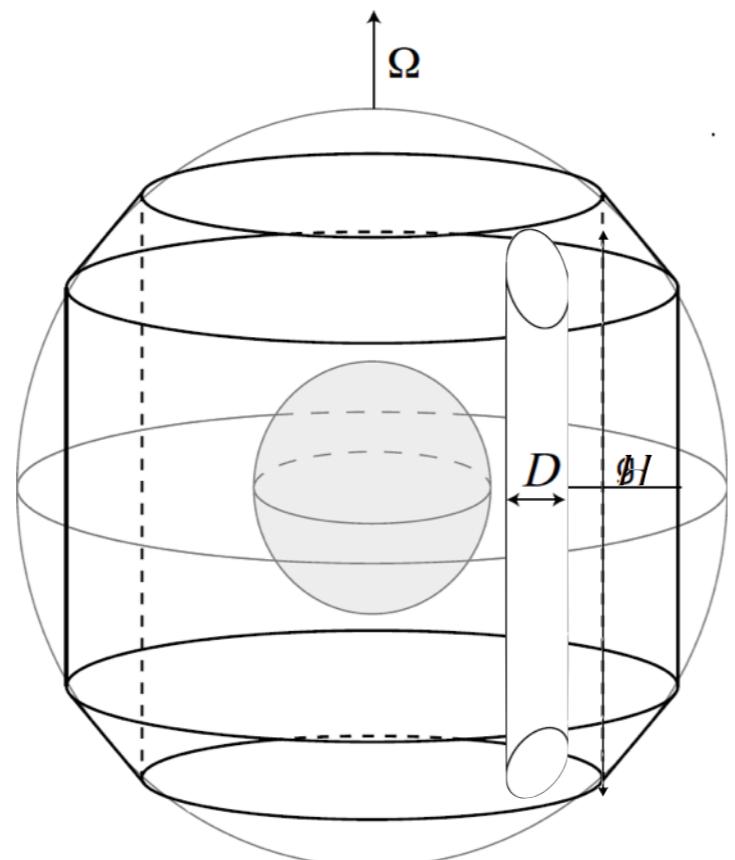


penetrative



J., Knobloch PoF 99, JMP '07 Sprague et al JFM 06, J. et al JFM 06 Groom et al PRL '10, J. et al GAFD 12, J. et al PRL 12

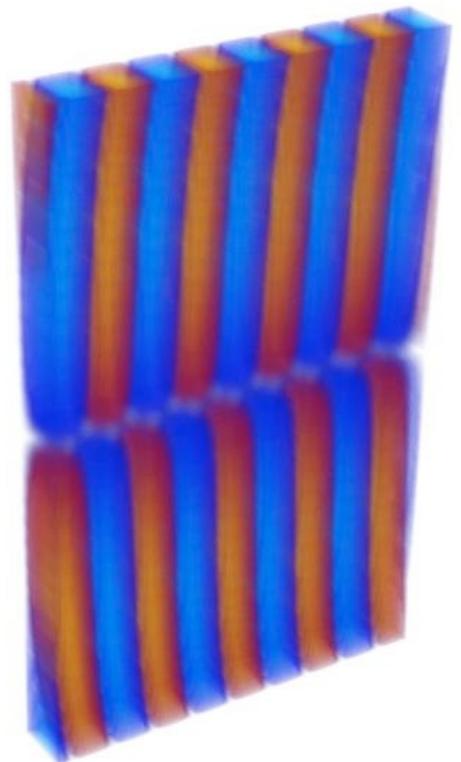
Planetary scale
convection:



Calkins et al JFM 13



thermal Rossby waves



QuasiGeostrophic Rayleigh-Bénard Convection

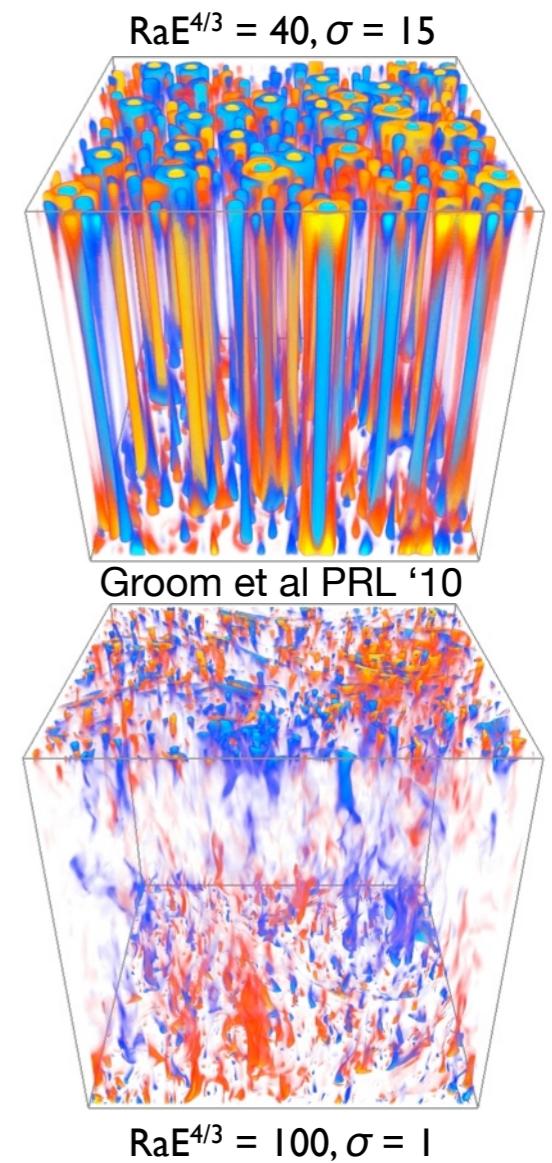
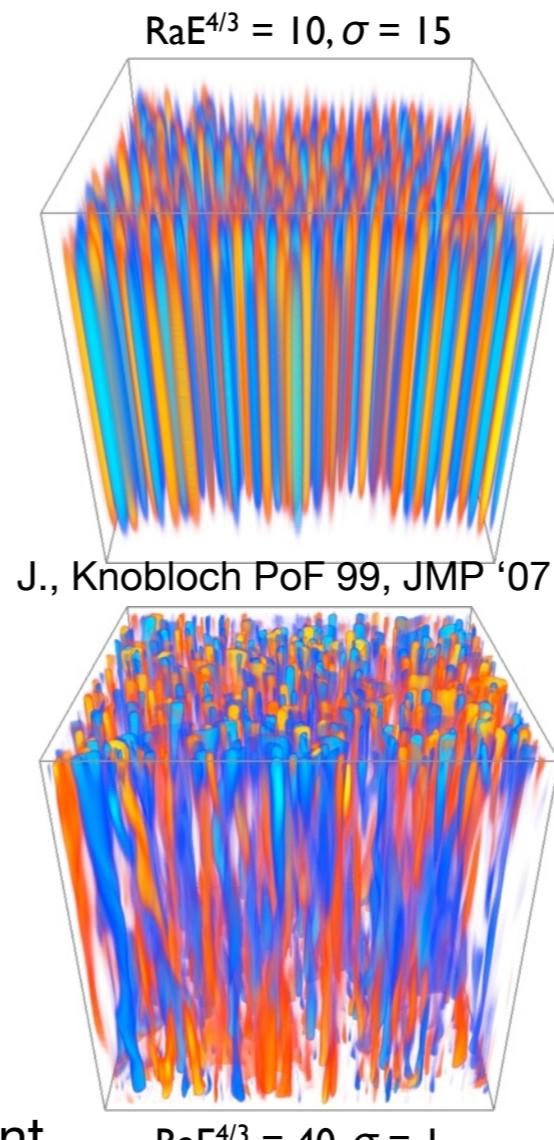
$$\partial_t \zeta + J[\psi, \zeta] - \partial_Z w = \nabla_{\perp}^2 \zeta$$

$$\partial_t w + J[\psi, w] + \partial_Z \psi = \nabla_{\perp}^2 w + \frac{RaE^{4/3}}{\sigma} \theta$$

$$\partial_t \theta + J[\psi, \theta] + w \partial_Z \bar{T} = \frac{1}{\sigma} \nabla_{\perp}^2 \theta$$

$$\partial_Z \bar{w\theta} = \frac{1}{\sigma} \partial_{ZZ} \bar{T}$$

- ▶ Four Flow Regimes as Ra \uparrow : laminar to turbulent



RaE^{4/3} = 100, σ = 1



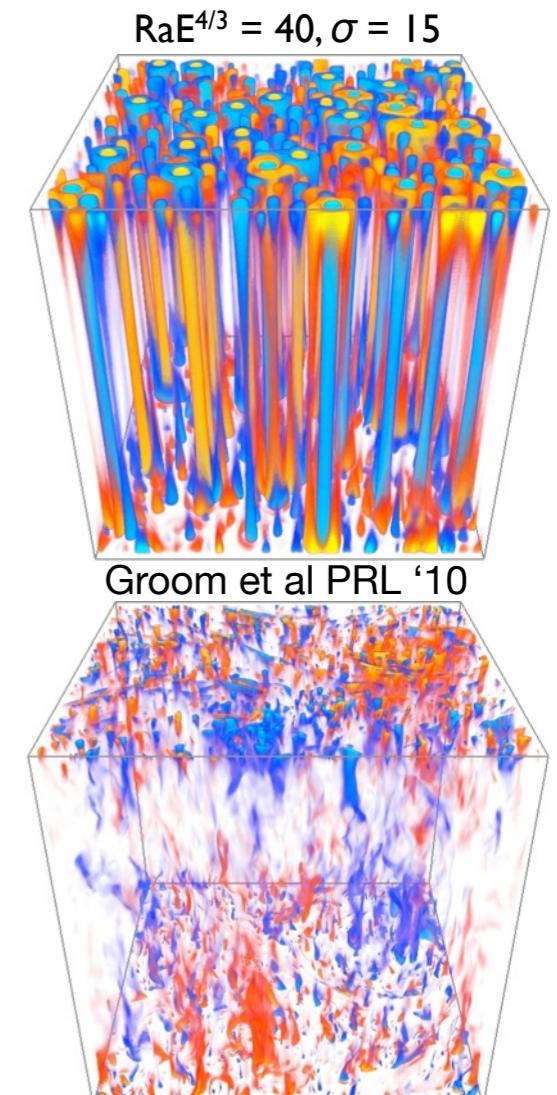
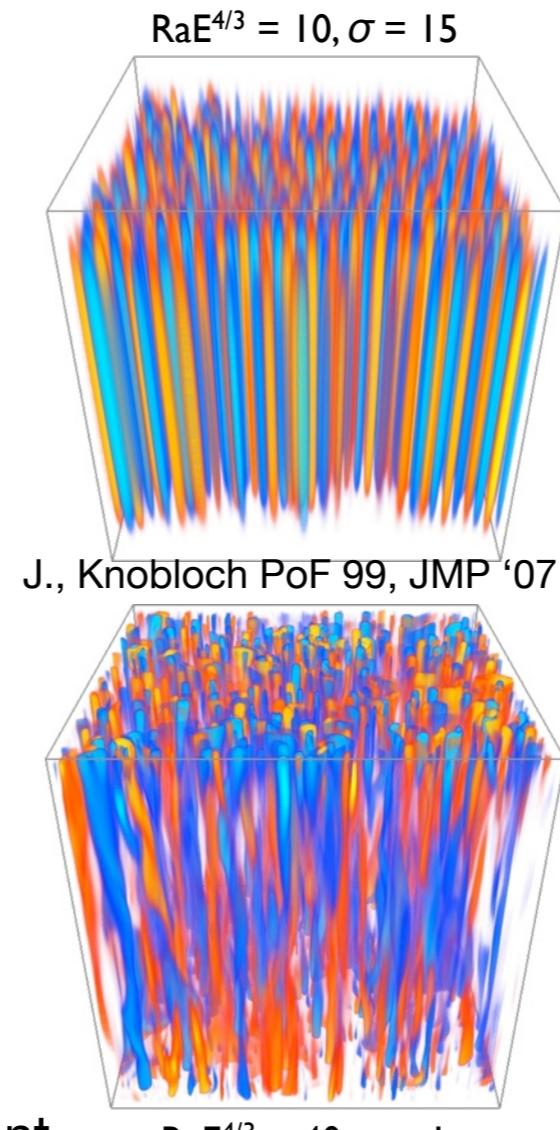
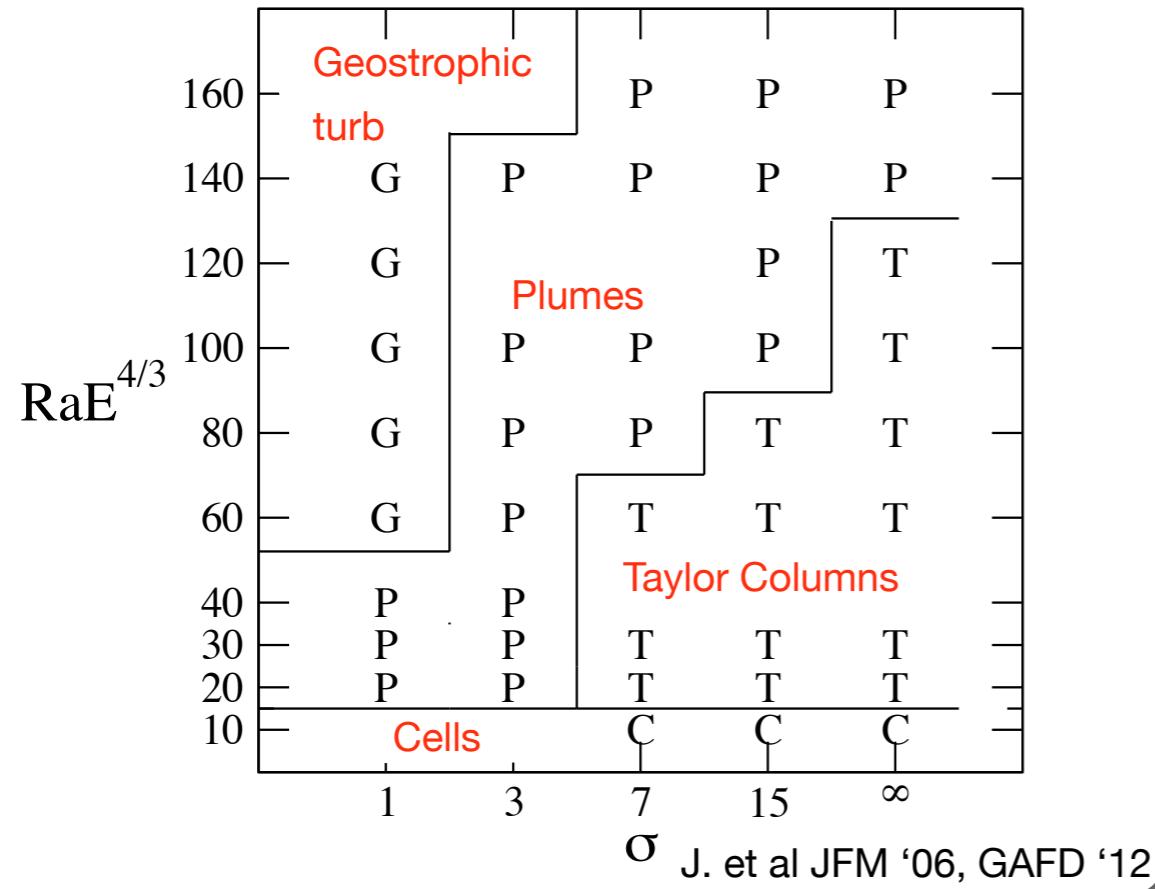
Cells \rightarrow CTC's via TBL instability & synchronization of TBL's

CTC's: Shielded vortical columns with zero circulation

Plumes regime occurs when TBL are unable to synchronize

- ▶ Ultimate Regime Geostrophic Turbulence (Julien et al GAFD 2012)

QuasiGeostrophic Rayleigh-Bénard Convection



Ω ↑
g ↓

► Four Flow Regimes as $\text{Ra} \uparrow$: laminar to turbulent

$\text{RaE}^{4/3} = 40, \sigma = 1$

$\text{RaE}^{4/3} = 100, \sigma = 1$

J. et al GAFD '12

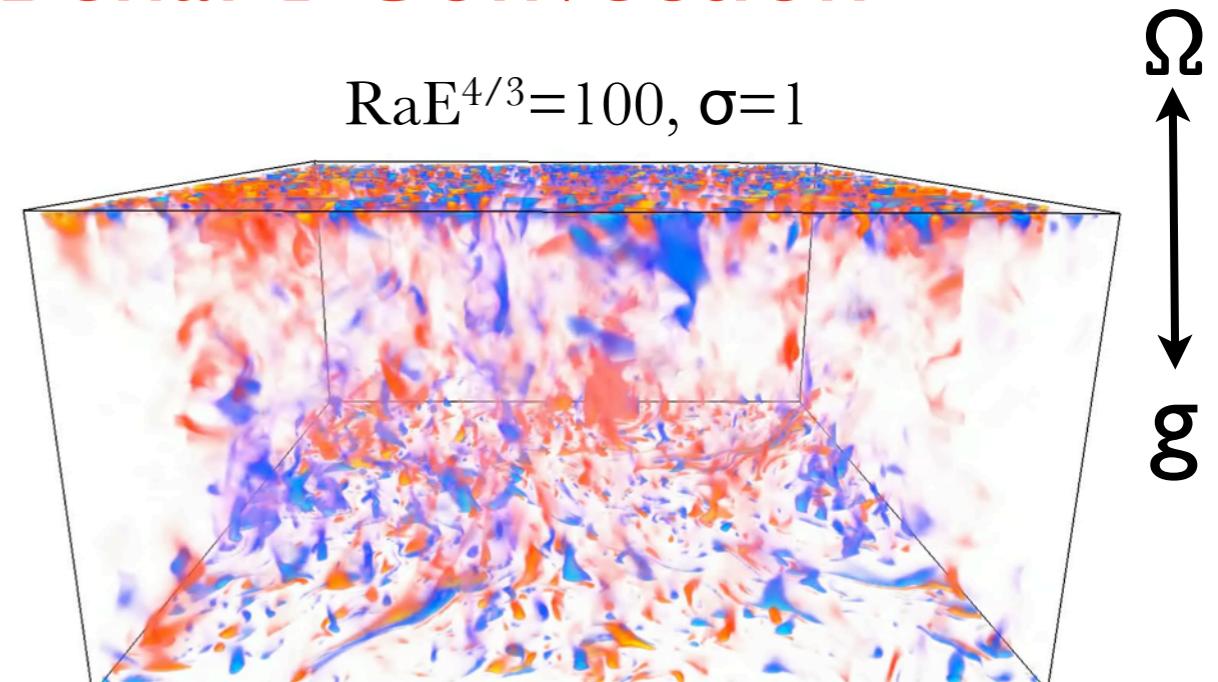
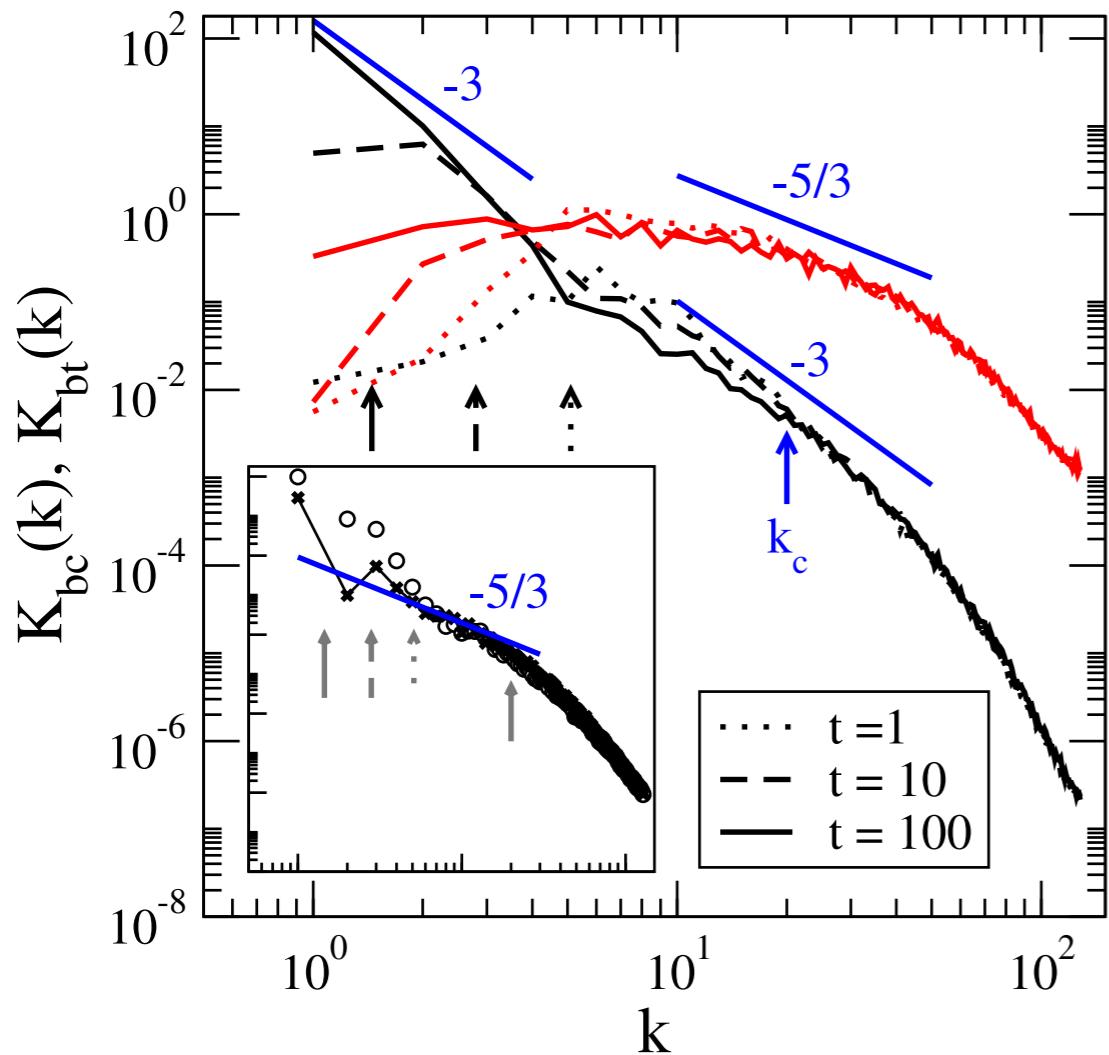
Cells → CTC's via TBL instability & synchronization of TBL's

CTC's: Shielded vortical columns with zero circulation

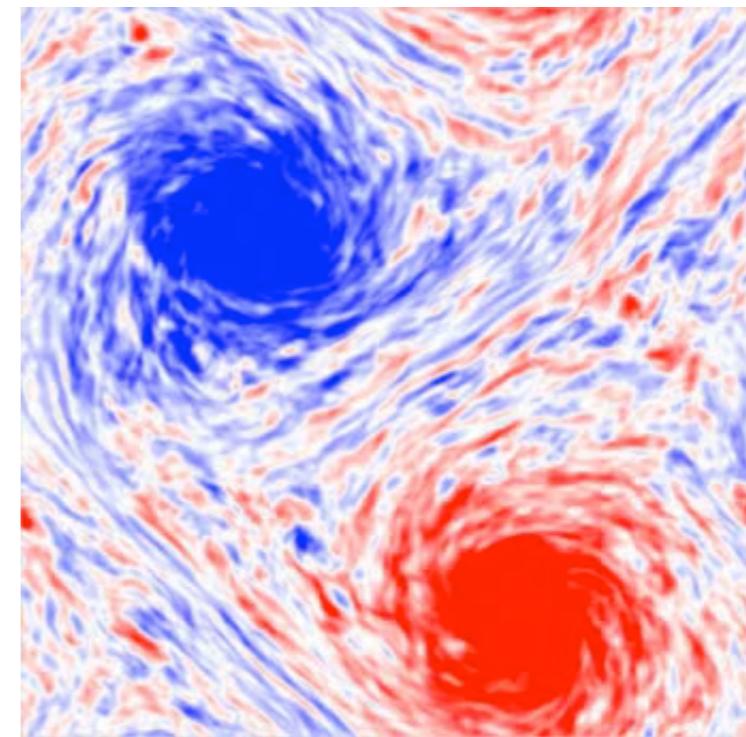
Plumes regime occurs when TBL are unable to synchronize

► Ultimate Regime Geostrophic Turbulence (Julien et al GAFD 2012)

Geostrophic Rayleigh-Bénard Convection



J. et al GAFD '12; Rubio, J., Weiss submitted '13



► Turbulent Inverse Cascade (J. et al GAFD '12, Rubio et al 2014)

Positive feedback loop

- GT provides the nonlinear forcing that generates BV's
- BV organizes GT thru advection and stretching
- BV produces large scale forcing to sustain itself

Depth averaged vorticity

► Energy spectra consistent with 2D barotropic and 3D baroclinic dynamics

Heat Transport by GT Convection

$$Nu - 1 = \frac{1}{25} \sigma^{-\frac{1}{2}} \left(RaE^{\frac{4}{3}} \right)^{\frac{3}{2}}$$

- Thermal throttling: region that controls the efficiency of heat transport in the fluid layer. TBL ? or Interior?

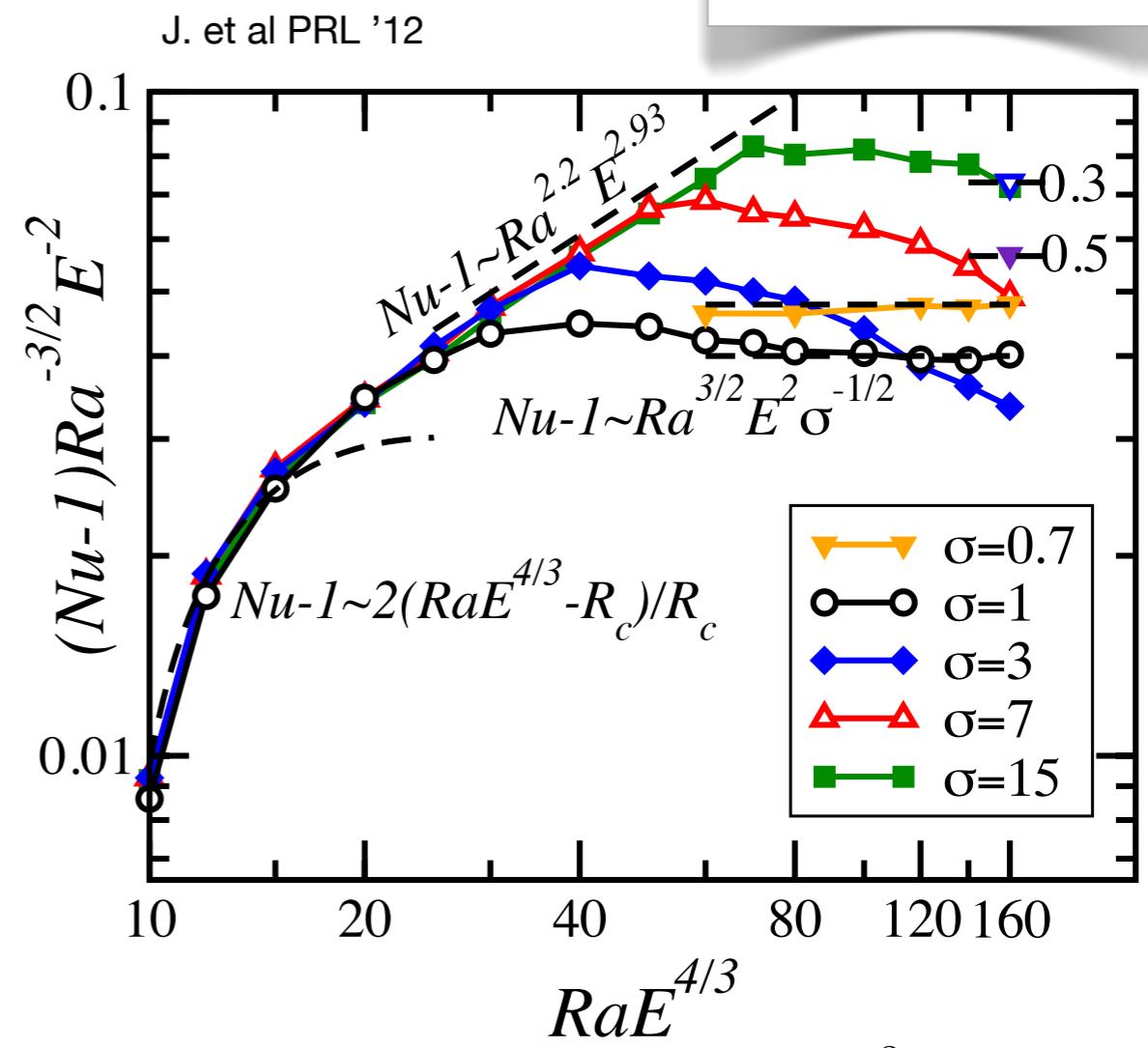
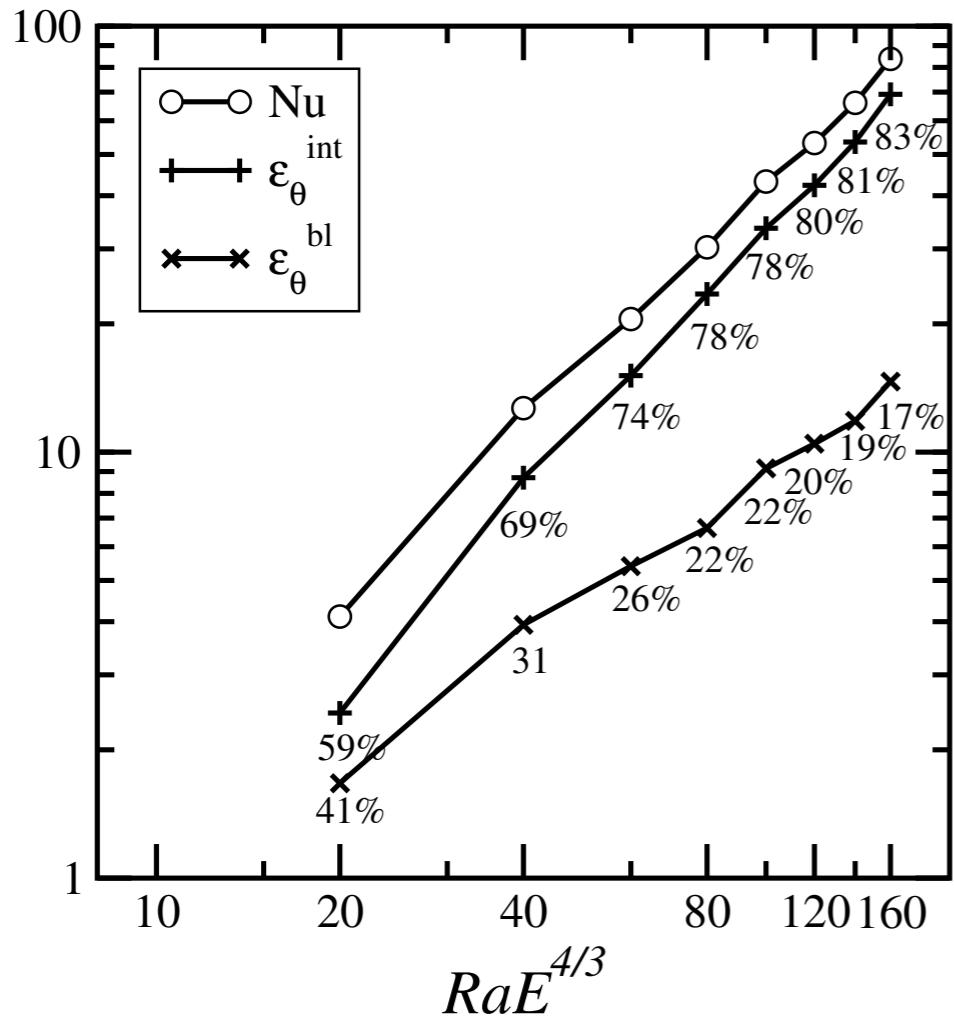
turbulent interior $\beta \Rightarrow$ dissipationless scaling law (Kraichnan '63, Howard '63)

TBL $\beta \Rightarrow$ marginally stably BL's (Malkus '63)

$$\beta_{turb} = 3/2$$

$$\beta_{tbl} = 3$$

rotating



$$\begin{aligned} \mathcal{E}_\theta &\approx \mathcal{E}_\theta^{int} = \langle |\partial_Z \bar{T}|^2 \rangle + \langle |\nabla_\perp \theta|^2 \rangle \\ &\equiv Nu \end{aligned}$$

$$\begin{aligned} Nu &\equiv \frac{QH}{\rho_0 c_p \kappa \Delta T}, \quad Ra = \frac{g \alpha \Delta T H^3}{\nu \kappa}, \quad E = \frac{\nu}{f H^2} \\ \sigma &= \frac{\nu}{\kappa} \end{aligned}$$

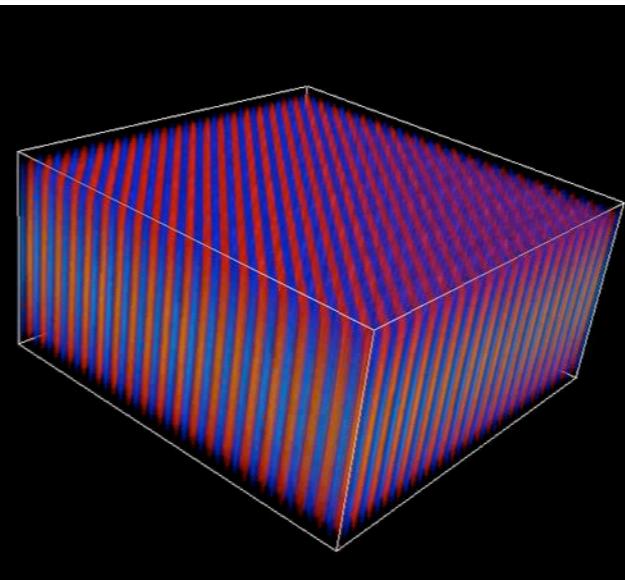
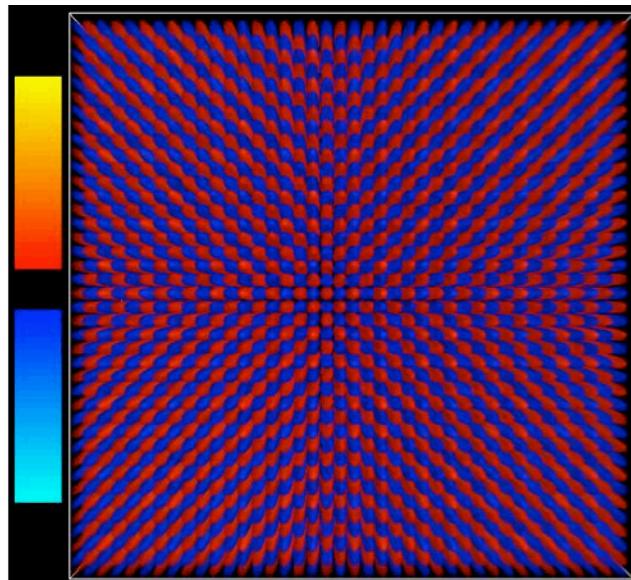
$$Ra_{crit} \propto E^{-4/3}$$

Outlook for 3D NH-QG

Thank you

- Reduced PDE's well suited to NHQG dynamics, computationally less challenging.
- Incompressible/Anelastic aDNS ("a"symptotic)
 - Investigate route to turbulence
 - Mean flow generation: inverse turbulent cascade?
 - Efficiency of heat transport: turbulent scaling laws
- ? Multiscale modeling: Coupling to balanced large-scale dynamics
Grooms, Fox-Kemper & J DAO '11
- Planetary convection: deep spherical shells

Sprague et al JFM '06 Groom et al PRL '10



Julien et al GAFD '12 Julien et al PRL '12

